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Review

An overview on SAR measurements of sea surface wind

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Abstract

Studies show that synthetic aperture radar (SAR) has the capability of providing high-resolution (sub-kilometer) sea surface wind fields. This is very useful for applications where knowledge of the sea surface wind at fine scales is crucial. This paper aims to review the latest work on sea surface wind field retrieval using SAR images. As shown, many different approaches have been developed for retrieving wind speed and wind direction. However, much more work will be required to fully exploit the SAR data for improving the retrieval accuracy of high-resolution winds and for producing wind products in an operational sense.

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Keywords: SAR; Sea surface wind; Wind retrieval algorithms

1. Introduction

Sea surface wind is a crucial parameter for the studies of sea surface variables such as waves, ocean circulation, marine meteorology and the coupling of oceanic and atmospheric systems. These variables can be measured from space by both active and passive microwave instruments such as scatterometers (SCATs), altimeters, SARs and radiometers. Among these instruments, satellite SCATs can measure global winds with a spatial resolution of up to 25 km on an operational basis. The SCAT winds, with the accuracy of wind speed and direction of $\pm 2 \text{ m s}^{-1}$ and $\pm 20^{\circ}$, respectively, have been assimilated into forecast models for a variety of weather and climate studies. However, the relatively low spatial resolution makes the SCAT winds best suited for the open ocean studies. Thus the observation of short-scale wind features from SCATs is not possible. In contrast, satellite SARs with all-weather capability and large spatial coverage, have a higher spatial resolution and therefore have the potential to provide wind

fields over the sea surface with spatial resolution nearly two orders higher (sub-kilometer) than conventional SCATs. This will enhance the understanding of marine atmospheric boundary layer and oceanographic processes, especially in coastal areas.

For incidence angles between 15° and 70° [1], the radar backscatter from the sea surface is primarily caused by the small-scale sea surface roughness, which is strongly influenced by the local wind field. This creates the possibility of extracting sea surface wind from SAR images. Early investigations of extracting the wind field from SAR images were made in Refs. [2] and [3], which demonstrated a correlation between the L-band SEASAT SAR image intensity and SEASAT scatterometer wind speed. Then Ref. [4] described the linear features aligned with the direction of the wind on the scale of a few kilometers in SEA-SAT images. Using Fourier transform techniques, the wind direction, albeit with a 180° ambiguity, was estimated from the wind row caused linear features. Following SEA-SAT, several satellites carrying SARs which work at different bands and polarizations have been launched. Based on the large amount of SAR data, different wind retrieval methods were then developed.

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In this paper, we will give a brief overview of the latest study on wind field retrieval from SAR images. The methods for estimating only of wind speed and wind direction are introduced in Sections 2 and 3, respectively. In Section 4 methods for simultaneous retrieval of wind speed and wind direction from SAR images are presented. Section 5 is conclusions.

2. Wind speed retrieval

Among the approaches for wind speed retrieval the most widely used ones are based on the methods of retrieving sea surface wind speed from SCATs. These methods are introduced in the first part. Some other alternative approaches for solely wind speed retrieval are presented in the second part of this section.

2.1. Scatterometry-based approaches

In scatterometry-based approaches, the dependency of the normalized radar cross-section (NRCS) of radar sensors (e.g., SCATs or SARs) on the wind speed and the geometry of observations (e.g., incidence angle and azimuth angle with respect to wind direction) is generally expressed by the geophysical model function (GMF). We summarize the general form of the function as

$$\sigma_0 = A(U_{10}, \theta) [1 + b_1(U_{10}, \theta) \cos \phi + b_2(U_{10}, \theta) \\ \times \cos 2\phi]^B$$
(1)

where σ_0 is the NRCS in linear units; ϕ is the relative wind direction, which represents the angle between the radar look direction and the local wind direction; U_{10} is the statistically neutral wind speed normalized to 10 m height; θ is the incidence angle and A, B, b_1 and b_2 are coefficients depending on U_{10} , θ , radar frequency and polarization.

Fig. 1 shows a general flow for wind retrieval from a SAR image using scatterometry-based approaches. With wind directions from either the SAR image itself or from

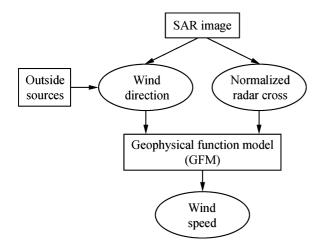


Fig. 1. Schematic description of wind retrieval from SAR images using scatterometry approaches.

outside sources (e.g., *in situ* data, SCAT measurements, or atmospheric model outputs) and a GMF, the wind speeds can be derived from the SAR image.

2.1.1. Wind speed retrieval from VV-polarized SAR

For the instruments operating at C-band and vertical (VV) polarization in transmission and reception, several empirical GMFs have been developed and validated through a series of satellite scatterometer missions. The most widely used are CMOD4 [5] developed and empirically tuned to ECMWF (the European Center for Medium-Range Weather Forecast) weather model results, CMOD_IFR2 [6] developed at Ifremer-France and calibrated against in situ measurements, in particular NOAA buoys and partly ECMWF model results, and CMOD5 [7] which was an upgrade of CMOD4 to address certain wind speed biases in the CMOD4 model function. It has been shown that these functions are applicable for wind speed retrieval from VV-polarized SARs onboard ERS-1 and ERS-2 and ENVISAT [8-14]. Fig. 2 shows the SARestimated NRCS as a function of sea surface wind speed. The NRCSs are computed by CMOD4, CMOD_IFR2 and CMOD5, respectively. Although CMOD5 is more suitable for estimating high wind speeds stronger than 25 m s⁻¹ [7]. Fig. 2 shows that all CMOD functions produce very similar results for wind speeds below 20 m s⁻¹. For different regions where GMF is the most appropriate, it is still an open question.

For L-band and VV-polarized SAR, a model function has been developed [15]. The L-band model enables wind

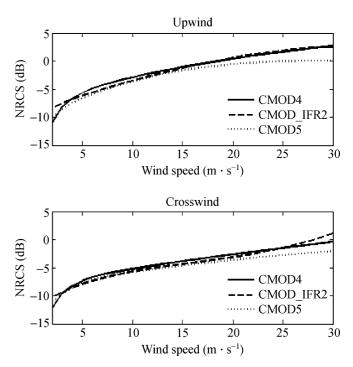


Fig. 2. NRCS as a function of the wind speed computed by various Cband model functions as indicated in the figure. The incidence angle is 25° , and the relative wind directions are 0° (upwind, upper panel) and 90° (crosswind, lower panel), respectively.

speed retrieval from SAR onboard JERS-1 (Japanese Earth Resources Satellite-1).

2.1.2. Wind speed retrieval from HH-polarized SAR

For wind speed retrieval from SAR images acquired at horizontal (HH) polarization in transmission and reception (e.g., SAR onboard RADARSAT-1), no well developed models exist, but two approaches have been attempted. The first and more widely used approach has been to develop a hybrid model function consisting one of the VV-polarized GMFs and a polarization ratio [16–22]. The second has been to derive analytical models directly for HH polarization [23].

In the first approach, the polarization ratio Pr is defined as

$$\Pr = \frac{\sigma_0^{VV}}{\sigma_0^{HH}} \tag{2}$$

where σ_0^{VV} and σ_0^{HH} are the NRCS values in VV and HH polarizations, respectively.

The polarization ratio has been measured by Unal et al. [24] with an airborne SCAT for incidence angles of 20°, 30° and 45° for wind speeds at 2–14 m s⁻¹. Their study showed that the ratio is mainly dependent on the incidence angle. But for wind speeds below 6 m s⁻¹, a wind speed dependency was also observed. By fitting a model to the data of Ref. [24], Thompson et al. proposed a polarization ratio model relating Pr to the incidence angle in the following form [25]:

$$\Pr = \frac{(1 + 2\tan^2\theta)^2}{(1 + \alpha\tan^2\theta)^2}$$
(3)

where Pr is in linear units; α is a constant relating to the type of surface scattering and was set to 0.6. This relation keeps a general form consistent with both the Bragg condition ($\alpha = 0$) and the geometrical optics (or Kirchhoff) condition ($\alpha = 2$). Ref. [26] compared RADARSAT SAR wind speeds derived by the CMOD2_IFR2 model with buoy measurements and determined that $\alpha \approx 1$, which was also suggested by Horstmann et al. [16]. From five dual-polarization (HH and VV) ENVISAT ASAR images, Ref. [27] expressed the polarization ratio as a polynomial function of tan(θ) for incidence angles below 30° and made the ratio follow the form of Eq. (3) with $\alpha = 1$ at incidence angles higher than 30°.

Another form of the polarization ratio is [28]

$$\Pr = \frac{(1+2\tan^2\theta)^2}{(1+2\sin^2\theta)^2}$$
(4)

which yields NRCS values in HH polarization in better agreement with the SASS-2 model derived from the Kuband SASS (Seasat-A Satellite Scatterometer) measurements than the classical one.

Based on the data from C-band airborne radar STORM, Ref. [29] proposed two analytical formulations to model the polarization ratio, one as an exponential function of the incidence angle only and the second one with additional dependency on the azimuth angle. The models have been assessed by ENVISAT ASAR images.

Fig. 3 shows the polarization ratio of (3) with α of 0.6 and 1, the form from Ref. [27], and the ratio of (4), together with the exponential form from Ref. [29]. The figure shows that the polarization ratios differ from each other with exception for those in Refs. [25] and [27] at incidence angles higher than 30°. The difference between various polarization ratios may be due to the sensitivity of the polarization ratio to the calibration uncertainty of different radars.

2.2. Alternative approaches

Besides the scatterometry-based approaches, several emerging methods for wind speed retrieval have also been developed in the recent years.

2.2.1. Azimuth cut-off method

The fundamental idea behind this method is the extraction of the wind speed from the sea surface wave spectrum. Based on the evaluation of 1200 SAR wave mode images with a central incidence angle of 20.2°, in the case of a fully developed sea (no fetch limitation), an empirical relation between U_{10} and the azimuth cut-off wave length λ_c was given in Ref. [30] as

$$U_{10} = 4.75 \left(\frac{\lambda_{\rm c} - 30}{110} \right) \tag{5}$$

The cut-off wave length can be estimated in the spectral domain either by fitting the auto covariance function (ACF) derived from an inverse Fourier transform of the SAR image power spectrum with a Gaussian function [31] or by fitting the ACF with more sophisticated models [32].

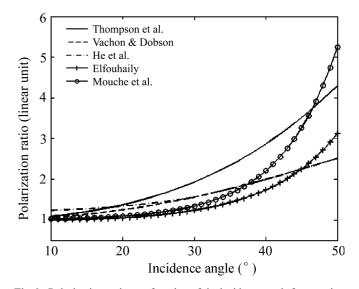


Fig. 3. Polarization ratio as a function of the incidence angle from various formulations as indicated in the figure.

The main advantage of this approach is that no information on wind direction is required. However, the approach only holds for C-band, VV polarization and low incidence angles ($\sim 20^\circ$). When using this approach, one must also be cautious since the cut-off wave length estimate includes long wave orbital motion contributions which may contaminate the result.

2.2.2. Neural network method

For uncalibrated SAR images, the neural network (NN) method allows the retrieval of wind speed directly from the SAR image intensity, independent of any knowledge of the NRCS. The method does not require explicit models for the SAR imaging process and can therefore be applied to any system configuration, i.e., polarization, incidence angle, etc. [33]. Ref. [33] applied this method to wind retrieval from SLC ERS-2 SAR wave mode images on a global and continuous basis, and the derived wind speeds agreed well with ERS-2 SCAT measurements. The NN method may also be applicable for retrieving both

wind speed and wind direction from calibrated images. However, no results from such investigations have been reported.

2.2.3. Physical model

Studies show that the contribution of specular reflection to the NRCS received by the radar is also significant at incidence angles less than 30° [34]. Considering the contribution of both specular reflection and Bragg scattering to NRCS, a physical model [35] was applied to wind speed retrieval from SAR images. The physical wind model can retrieve wind speeds from either VV- or HH-polarized SAR images. The comparison of wind speeds measured by five buoys in Hong Kong coastal waters to those retrieved from 30 VV-polarized ENVISAT ASAR images from October 2005 to June 2007 using the CMOD4, CMO-D_IFR2, CMOD5 and the physical model, respectively, shows that the physical model is comparable with the Cband empirical models (see Fig. 4a-d where RMSE, Bias and R denote root mean square error, bias and correlation coefficient, respectively).

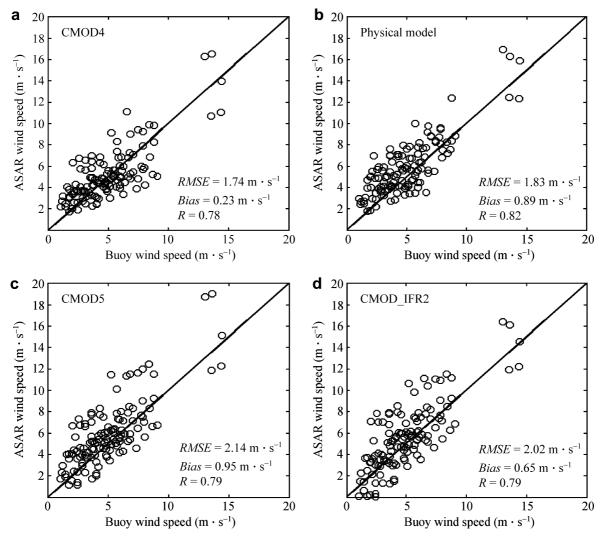


Fig. 4. Comparison of buoy measurements with wind speeds retrieved from VV-polarized ENVISAT ASAR images using CMOD4 (a), physical model (b), CMOD5 (c) and CMOD_IFR2 (d), respectively.

3. Wind direction retrieval

To measure wind speeds from SAR images using wind direction dependent methods such as the scatterometrybased approaches, the main difficulty is the lack of knowledge of the wind direction. To solve this problem, there are different methods in use, for example, assuming a fixed direction from a measurement for the whole image, or interpolation of the wind direction from SCATs or atmospheric models. Despite the empirical success of SAR wind speed retrieval using such wind direction, the above approaches ignore wind direction information in the SAR image itself. In SAR images, if the linear features aligned with the wind direction are visible, the wind direction can be derived directly from the images. These features, however, could be caused by any of the several physical effects, e.g., marine atmospheric boundary layer rolls [4,36], Langmuir cells, surfactant streaks, or wind shadowing [8,37,38].

Several methods have been developed to retrieve the orientation of the linear features aligned with the wind direction in SAR images. These methods can be divided into two categories: one is in the spectral domain, the other in the spatial domain.

In the spectral domain, the most popular method is the Fast Fourier Transform (FFT) method, which was first described by Gerling [4] and then modified and applied by several others [8,10,16,17,26,31,37,39-41]. In this method, the Fourier spectrum of the SAR image is computed. The main spectral energy is located perpendicular to the orientation of the wind streaks, giving a wind direction with a 180° directional ambiguity. The 180° directional ambiguity is a very important issue and is hard to remove. It can be removed only if wind shadowing is present which is often visible in the lee of coastlines, or if there are additional data available, e.g., wind direction from in situ data, SCAT measurements, or atmospheric model outputs. With a relatively low-resolution of wind direction retrieval such as 20×20 km², the FFT method works fine on open oceans.

A significant problem with the FFT method is that there may not be a feature aligned with the wind in all the parts of the image. Thus the method will often generate erroneous directions over featureless regions of the image. To address this problem, Ref. [42] developed a projection method based on projections of the image values at various angles, which can automatically estimate wind direction only from the image locations containing sufficient linear features. For the regions with no appropriate features, wind directions can be interpolated from the surrounding directions.

In the spatial domain, a recently developed wind direction retrieval method is the local gradient (LG) method, which derives the orientation of the wind streak by computing the local gradient on different scales [12,18,43]. Alternative approaches such as wavelet analysis [44,45] and variance method [46] have also been applied. Compared with the spectral analysis, the spatial analysis improves the resolution of SAR direction retrieval to better than 10×10 km² and even to 1×1 km² [43].

Although wind direction can be derived from the linear features on SAR image itself, it should be noted that not all SAR images contain linear features aligned with the wind direction. Moreover, the presence of long wind waves and other atmospheric features visible in some images are difficult to separate from these linear features. To solve these problems, emerging methods have been developed for retrieving wind direction from SAR images. These methods, to be introduced in Section 4, are capable of estimating wind speed simultaneously and have been called wind vector retrieval methods.

4. Wind vector retrieval

4.1. Optimal inversion method

Based on a general statistical approach to solve inversion problems (including underdetermined problems) in meteorological analysis, Ref. [47] proposed a SAR wind retrieval method which can estimate wind speed and direction simultaneously. This simplified method combines the SAR information, i.e., NRCS and/or wind induced streaks, with some background information coming from high-resolution limited area model or buoy data to retrieve the most probable wind speed and wind direction, assuming that all sources of information contain errors and that these are well characterized. The comparison of the wind vectors retrieved from this method and that from the scatterometry approach indicated that the optimal inversion method showed a promising result. However, a priori information (background wind data) would be required for this approach.

4.2. Multiple NRCS method

This method was recently developed by He et al. [48] and was applied to ASAR images. It was motivated by the methodology used for SCAT data, in which the wind vectors are determined from multiple NRCSs measured by SCATs with different look directions. In this method, a given SAR image was divided into many blocks along the radar look direction and the radar flight direction, respectively. Wind vectors can be retrieved by using multiple NRCSs from two neighboring blocks along the radar look direction, which have slightly different incidence angles. A constraint on this method is that the wind speeds retrieved from ASAR images in the near range tend to be underestimated, which may be due to the underestimation of NRCS related to the saturation effect of the SAR analog-to-digital conversion (ADC) in the near range.

4.3. Dual-polarization method

The method for retrieving the wind vector using dualpolarization ASAR images was developed by Song et al. [49]. Based on the combination of two ASAR images with co-polarization and cross-polarization, the wind vector can be retrieved by solving three equations with only two

Table 1
Summary of approaches for sea surface wind retrieval

Wind speed	Wind direction	Advantages	Disadvantages
Scatterometry-based approaches	/	Validated by large amounts of data. In quasi-operational use	One-dimensional wind component. Wind direction dependent
Azimuth cut-off method	/	Independent of wind direction and image calibration	One-dimensional wind component. Limited to low incidence angle. Fetch dependent
Neural network method	/	Independent of image calibration. Available for any system configuration (radar frequency, incidence angle, polarization, etc.)	One-dimensional wind component
Physical model	/	Available for any polarization	One-dimensional wind component. Wind direction dependent
/	Model output	Always available. Reasonable directions	Low-resolution. Space/time mismatch with SAR images
/	Alternative data (<i>in situ</i> , SCAT, etc.)	Independently validated directions	Low-resolution. Space/time mismatch
/	From linear features in SAR image (FFT, LG, etc.)	Space/time exact match. Higher- resolution than model output or alternative data	Not always available. Contaminated by other features and 180° ambiguity of wind direction
Optimal inversion method		Using both SAR data and background model data	Requires <i>a priori</i> information and a set of assumptions
Multiple NRCS method		Independent of <i>a priori</i> information	Underestimation of wind speed in the near range. 180° ambiguity of wind direction
Dual-polarization method		Independent of <i>a priori</i> information Eliminates 180° ambiguity of wind direction	Requires SAR with both co-polarization and cross-polarization

unknown variables, e.g., wind speed and wind direction. In this way, the 180° directional ambiguity occurring when a single SAR image is used can be effectively eliminated. But this method is only applicable for wind retrieval from SAR with both co-polarization and cross-polarization, such as ASAR.

5. Conclusions

SARs can measure sea surface wind fields with high-resolution at sub-kilometer scales. As mentioned above, different approaches have been developed to estimate wind parameters (wind speed and/or wind direction) from SAR images. Each approach has its advantages and drawbacks. Table 1 is a summary of these approaches. Among the different approaches, the scatterometry-based approaches for wind speed retrieval and FFT or LG methods for wind direction retrieval are the ones mostly used and have been applied to the quasi-operational environment [19]. Comparisons with buoy measurements in the previous studies show an accuracy of better than 1.8 m s^{-1} for wind speed in the 2–20 m s⁻¹ wind speed range and better than 20° for wind direction [20,38,48] with a 180° ambiguity. Comparisons with SCAT measurements show a random error of wind speed and wind direction even better than 1.0 m s^{-1} and 15° , respectively [21,33,49,50].

Although various wind retrieval approaches have been suggested, the sea surface wind retrieval from SAR images is still in its infancy, since it was developed only 10 years ago. Much more work is required to understand how we can fully exploit the SAR data for improving the retrieval accuracy of high-resolution winds. As data sets, including SAR images, sea surface winds and wind data from other sources, e.g., SCATs and atmospheric model outputs continue to grow, SAR wind retrieval performance can be statistically evaluated, high-resolution SAR winds can be provided in an operational sense, and the full evaluation of SAR winds can be realized.

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